

Optimal surface salinity perturbations of the Meridional Overturning Circulation

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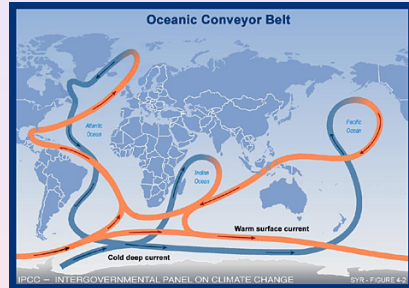


Climate context (1) : The MOC

Slow dynamics of the ocean :

Meridional Overturning Circulation
(MOC)

- Intensity of ~ 18 Sv
- Time scale of ~ 500 ans



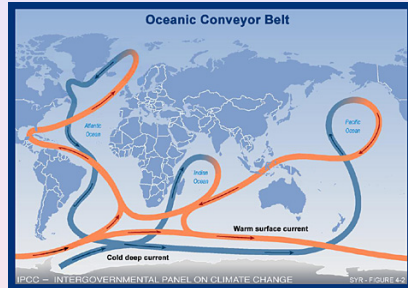
Northward transport of heat influencing the European climate
 \Rightarrow Variability of the meridional overturning circulation

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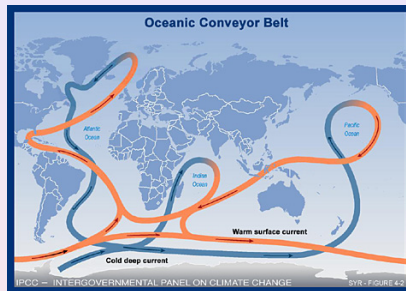
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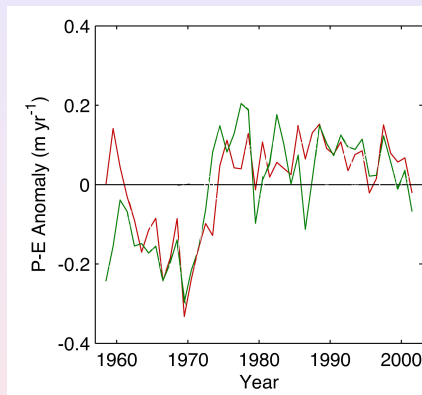


Northward transport of heat influencing the European climate

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Climate context (2) : North Atl. P-E

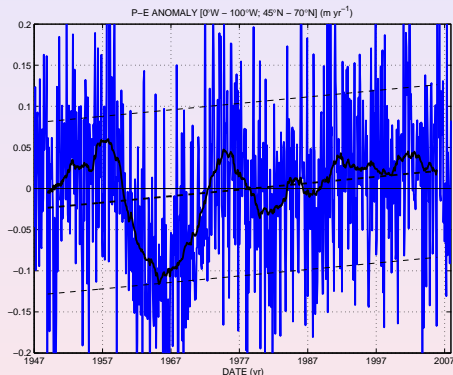
- Increase of precipitation in the north Atlantic
→ Josey and Marsh (2005)



What is the impact of the SSS modification on the meridional overturning circulation ?

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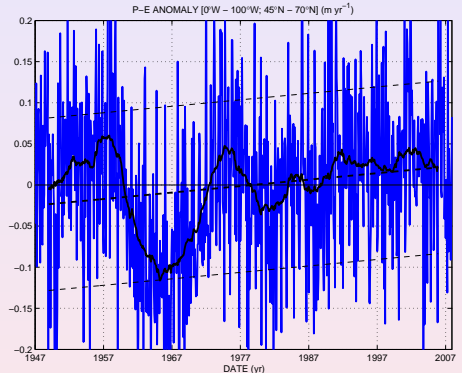
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Approach

⇒ Forced variability of the ocean circulation

- **Linear approach** :
weak variations (perturbations) of the ocean circulation
- Generalized stability analysis :
 - Atmosphere ⇒ optimal initial and stochastic perturbation (Farrell and Ioannou, 1996)
 - Ocean ⇒ optimal initial perturbation (Moore and Farrell, 1993)
 - 3 box THC ⇒ optimal initial and stochastic perturbation (Tziperman and Ioannou, 2002)

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Maximization method : Oceanic circulation application

Goal :

- Optimal impact of the SSS on the circulation

Lagrange parameters method

- Functions to maximize $\langle F|u(t) \rangle$ (or $\langle F|u(t) \rangle^2$) :
 - Meridional Overturning Circulation (MOC)
at the latitude and depth of its steady state maximum
(or its variance)
- Constraints
 - 1 Normalisation : $\langle u(0)|S|u(0) \rangle = 1$
 - 2 Salt conservation : $\langle C|u(0) \rangle = 0$
 - 3 Only surface salinity perturbation : $|u(0) \rangle = P|u' \rangle$

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Optimal perturbation experiments

	2D	PG	OPA - ORCA2
Initial perturbation	X	X	X
Constant perturbation	X		
Stochastic perturbation	X	X	

- Latitude-depth model
 ⇒ Methodological study
 Sévellec et al. (*J. Phys. Oceanogr.*, 2007)
- Planetary geostrophic model :
 ⇒ Influence of the surface boundary condition (flux vs mixed)
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Upper bound in the 2D model

- **Initial SSS perturbation :**
Great Salinity Anomalies (GSA, Belkin et al., 1998)
0.5 psu on 250 m \Rightarrow 2 Sv
- **Constant FW perturbation :**
hydrological cycle modification in the global warming scenario
(Held and Soden, 2006)
4% (3 cm yr⁻¹) \Rightarrow 0.14 Sv
- **Stochastic FW perturbation :**
Using 2 different reanalysis datasets
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- **Variability :**
 - Large scale gradient SSS **efficiently stimulates** a North Atl. multidecadal oscillation.
- **Surface boundary condition :**
 - The sensitivity pattern weakly depends of the surface boundary condition.
 - The intensity of the response strongly depends of the boundary condition.

Results in the PG model

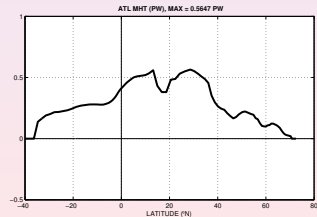
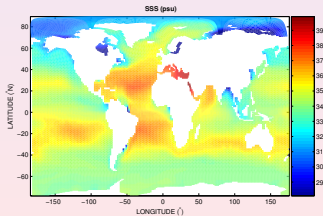
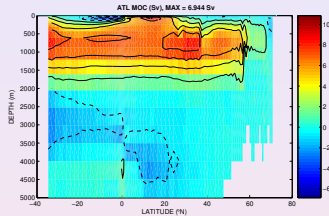
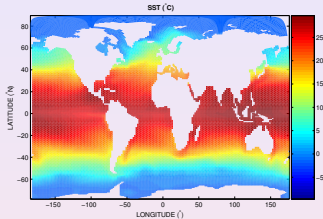
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Application in an Ocean General Circulation Model

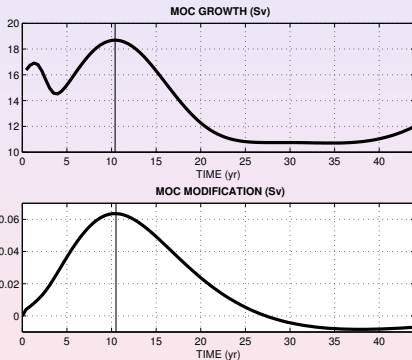
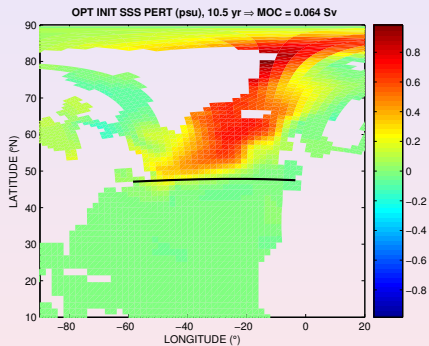
Model : OPA 8.2, ORCA2, OPATAM



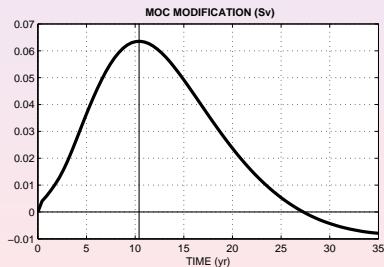
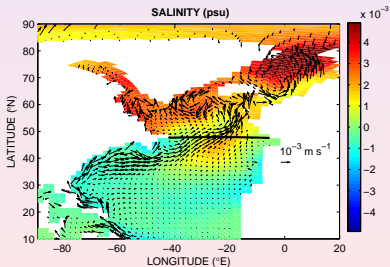
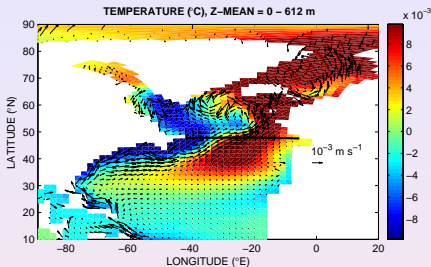
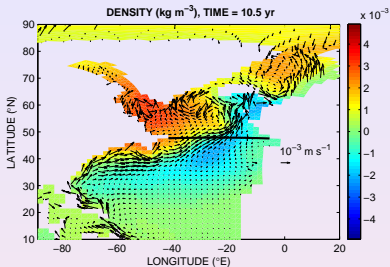
● MAX(MOC)=7 Sv (48°N)

● MAX(MHT)=0.6 PW (27°N)

Optimal initial SSS perturbation for the MOC



- Maximum growth after 10.5 yr

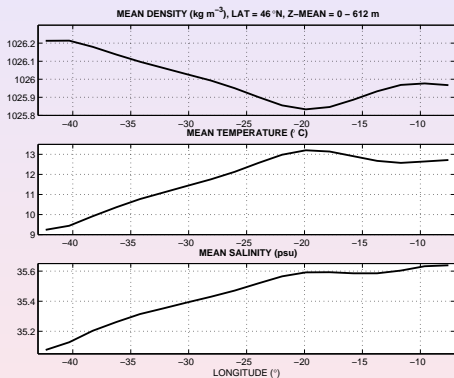


• $v'_{surf} > 0$

• $|\beta \partial_{\theta} S'| \gg |\alpha \partial_{\theta} T'|$

• $|\alpha \partial_{\phi} T'| \gg |\beta \partial_{\phi} S'|$

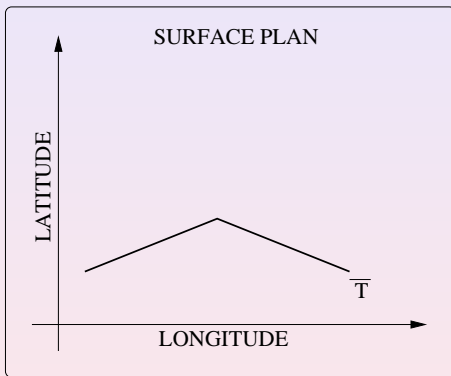
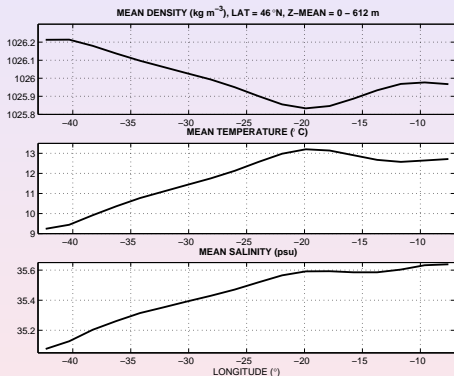
Finite time growth mechanism



$$\alpha \partial_{\phi} \bar{T} \gg \beta \partial_{\phi} \bar{S}$$

$$SSS'_{\text{north}} > 0 \Rightarrow v'_{\text{surf}} > 0$$

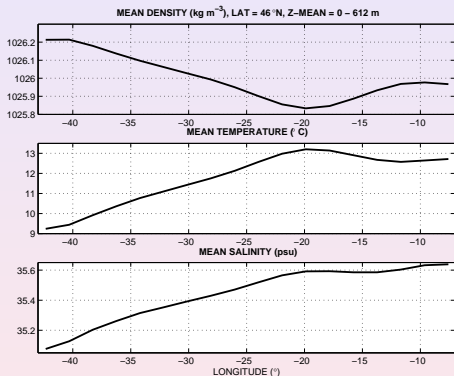
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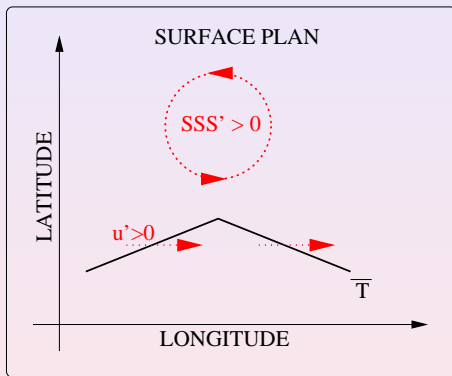
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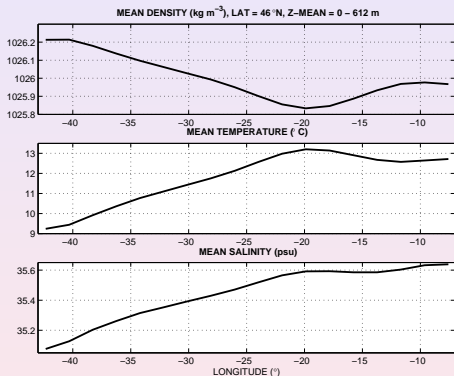


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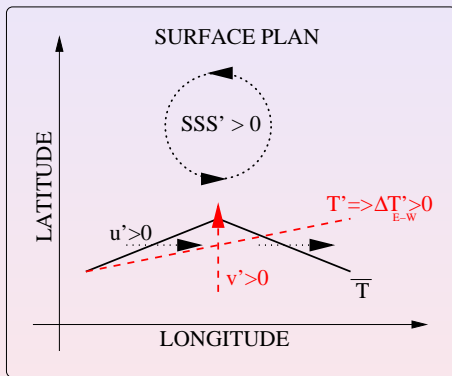


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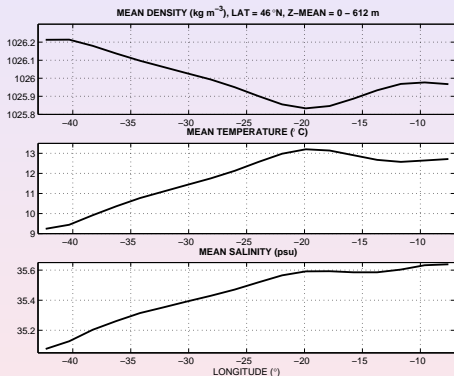


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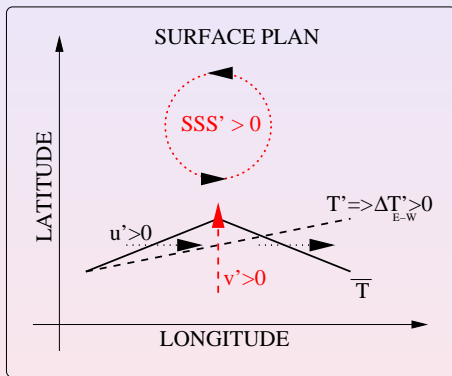


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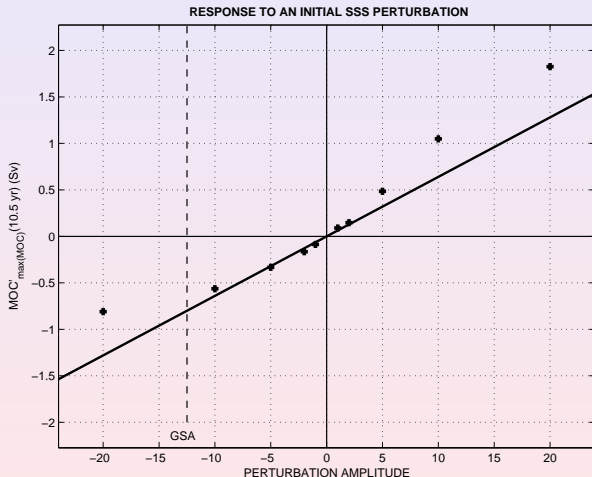


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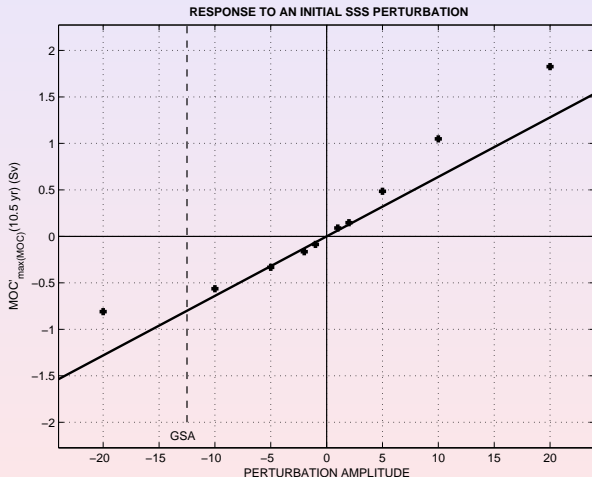
Nonlinear - linear comparison



- Relative error :
less than 20%

- Max bound :
GSA \Rightarrow 0.75 Sv
(11% of $\overline{\text{MOC}}$)

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- **Efficient method** to obtain the optimal initial perturbation :
⇒ **Explicit solution** (adj. model)
- Results of the 2D, PG and OGCM models
⇒ Similarity :
 - In the 2D, the similarity is controlled by the salinity
and the response is dominated by the Temperature
⇒ Difference : Transient growth mechanism
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Future work

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 - Impact of the Southern Ocean
 - Mechanism of the finite time growth
- Seasonal cycle (non-autonomous operator)
 - Sensitivity to the season
- Tropical study :
 - Optimal ocean perturbation and phase locking of ENSO (ENSEMBLES, European project for climate changes prediction)



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Thank you for your attention

Optimal initial SSS perturbation

Perturbation evolution (autonomous problem) :

$$\begin{aligned}\partial_t |u\rangle &= \mathbf{A} |u\rangle, \\ \Rightarrow |u(\tau)\rangle &= \mathbf{M}(\tau) |u(0)\rangle = e^{\mathbf{A}\tau} |u(0)\rangle.\end{aligned}$$

Explicit solution (using the adjoint model) of the optimal initial perturbation :

$$\begin{aligned}\Rightarrow |u(0)\rangle &= \mathbf{P} |u'\rangle \\ |u'\rangle &= (2\gamma_1)^{-1} \left(\mathbf{N}^{-1} \mathbf{P}^\dagger \mathbf{M}^\dagger(\tau) |F\rangle - \gamma_2 \mathbf{N}^{-1} \mathbf{P}^\dagger |C\rangle \right), \text{ with } \mathbf{N} = \mathbf{P}^\dagger \mathbf{S} \mathbf{P}, \\ \gamma_1 &= \text{fct} \left(\mathbf{M}^\dagger(\tau) |F\rangle, |C\rangle, \mathbf{N}, \mathbf{P}, \gamma_2 \right) \text{ and} \\ \gamma_2 &= \text{fct} \left(\mathbf{M}^\dagger(\tau) |F\rangle, |C\rangle, \mathbf{N}, \mathbf{P} \right).\end{aligned}$$

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\Rightarrow Solution depends on the maximization delay τ

Efficient method :

Maximization under constraints : $dG(\gamma, |u_0\rangle) = 0$

- Measure : **Linear function**

$$G(\gamma, |u_0\rangle) = \langle F | \mathbf{M}(\tau) | u_0 \rangle - \gamma (\langle u_0 | \mathbf{S} | u_0 \rangle - 1)$$

Explicit solution :

$$|u_0\rangle = \pm \frac{\mathbf{S}^{-1} \mathbf{M}^\dagger(\tau) | F \rangle}{\sqrt{\langle F | \mathbf{M}(\tau) \mathbf{S}^{-1} \mathbf{M}^\dagger(\tau) | F \rangle}}$$

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Eigenvalue solution :

$$\gamma \mathbf{S}_1 |u_0\rangle = \mathbf{M}^\dagger(\tau) \mathbf{S}_2 \mathbf{M}(\tau) |u_0\rangle, \quad \langle u_0 | \mathbf{S}_1 | u_0 \rangle = 1$$

Efficient method :

Maximization under constraints : $dG(\gamma, |u_0\rangle) = 0$

- Measure : **Linear function**

$$G(\gamma, |u_0\rangle) = \langle F | \mathbf{M}(\tau) | u_0 \rangle - \gamma (\langle u_0 | \mathbf{S} | u_0 \rangle - 1)$$

Explicit solution :

$$|u_0\rangle = \pm \frac{\mathbf{S}^{-1} \mathbf{M}^\dagger(\tau) | F \rangle}{\sqrt{\langle F | \mathbf{M}(\tau) \mathbf{S}^{-1} \mathbf{M}^\dagger(\tau) | F \rangle}}$$

- Measure : **quadratic norm**

$$G(\gamma, |u_0\rangle) = \langle u_0 | \mathbf{M}^\dagger(\tau) \mathbf{S}_2 | \mathbf{M}(\tau) | u_0 \rangle - \gamma (\langle u_0 | \mathbf{S}_1 | u_0 \rangle - 1)$$

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